Scientific basis of knowledge on Earth's composition

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I present, from a historical perspective, a logical progression of understanding related to the composition of the deep interior of the Earth that comes from fundamental discoveries and from discoveries of fundamental quantitative relationships in nature. By following step by step the reasoning from that understanding, one might begin to appreciate what is not yet known that pertains to recent interest in georeactor-produced antineutrinos and also what should be investigated to further advance that understanding.

Recent interest in detection of geo-antineutrinos with spectral and directional resolution¹⁻³ is welcomed as a potential means of verifying the existence of a nuclear georeactor at the center of the Earth⁴⁻⁶. Generally, as physicists venture into a new area, there is a learning curve and there may be some confusion. But as physicists venture into the area of solid-Earth science, confusion may be considerably magnified by an Earth science literature turgid with decades of reports of investigations that fail to follow long-established, ethical standards of science. One major problem is that many investigators make models, based upon arbitrary assumptions, or based upon other models, themselves based upon arbitrary assumptions. Such models, sometimes designated by the makers as 'reference model' or 'preferred model', often ignore contradictory scientific evidence and can generally be replaced by different models based upon other assumptions, all of which may have questionable relevance, utility, and correctness.

The purpose of science is not to make arbitrary models based upon assumptions, but, rather, to determine the true nature of the Earth and the Cosmos, which can be done by making fundamental discoveries and by discovering fundamental quantitative relationships in nature. In the following, I present from a historical perspective a logical progression of understanding about the Earth that comes from such fundamental discoveries and fundamental relationships. By following step by step the reasoning from that understanding, one might begin to appreciate what is not yet known that pertains to recent interest in geo-antineutrinos and also what should be investigated to further advance that understanding.

Historical basis

In 1897 Wiechert⁷ realized from the density determination of Cavendish⁸ that the

Earth cannot consist wholly of rock. Having seen in museums meteorites that consist entirely of nickeliferous iron metal, as well as meteorites made of both iron metal and stone, Wiechert suggested that the Earth has at its center a core, like the metallic iron of meteorites. The existence of such a core, Wiechert estimated, could account for the high bulk density of the Earth.

Oldham⁹ determined the speed of earthquake-waves as a function of depth of travel within the Earth. He found generally that beneath the crust the velocities of earthquake-waves increase with increasing depth, but only to a particular depth, below which their velocities become abruptly and significantly slower. Oldham had discovered the Earth's core.

During the next twenty-five years, the dimension of the Earth's core was determined precisely and its state was shown to be liquid due to its failure to support transverse earthquake-waves¹⁰. Seismological data, augmented with moment of inertia considerations, however, can yield information about dimensions, physical states, and mass distributions of structures within the Earth. But for elemental compositions one must rely upon implications derived from meteorites.

The constancy in isotopic compositions of most of the elements of the Earth, the Moon, and the meteorites indicates formation from primordial matter of common origin. Primordial elemental composition is yet manifest and determinable to a great extent in the photosphere of the Sun. The less volatile rock-forming elements, present in the outer regions of the Sun, occur in nearly the same relative proportions as in chondritic meteorites. But chondrites differ from one another in their respective proportions of major elements 11,12, in their states of oxidation^{13,14}, mineral assemblages¹⁵, and oxygen isotopic compositions¹⁶ and, accordingly, are grouped into three distinct classes: enstatite, carbonaceous and ordinary. Virtually all approaches to whole-Earth composition are based upon the idea that the Earth is similar in composition to a chondrite meteorite. A major problem within the Earth sciences began more than six decades ago with a wrong choice of chondrite type as being representative of the Earth.

When earthquake-waves enter and leave the core, they change speed and direction. Consequently, there is a region, called the shadow zone, where earthquake-waves should not be detectable. But in the early 1930s earthquake-waves were in fact detected in the shadow zone. Lehmann¹⁷ discovered the inner core by showing that a small solid object, within the fluid core, could cause earthquake-waves to be reflected into the shadow zone.

The contradiction

Four years later, Birch¹⁸ pronounced the composition of the inner core to be partially crystallized nickel-iron metal. It is important to understand Birch's logic, which, within the framework of knowledge at the time, seemed reasonable; but which ultimately set the Earth science community along an incorrect progression of development.

Birch envisioned the Earth to be like an *ordinary* chondrite meteorite, the most common type of meteorite observed to fall to Earth. He ignored the rare, oxygenrich *carbonaceous* chondrites, which contain little or no iron metal, and he ignored the rare oxygen-poor *enstatite* chondrites, which contain some minerals, such as oldhamite (CaS), that are not found in the surface regions of the Earth.

Birch thought that nickel and iron were always alloyed in meteorites and he knew that the total mass of all elements heavier than nickel was too little to comprise a mass as large as the inner core. Birch therefore assumed that the inner core was nickel—iron metal that had begun to crystallize from the melt. That assump-

tion, which underlies much geophysical and geochemical development over the past six decades, is unfounded.

From discoveries made in the 1960s, I realized a different possibility for the composition of the Earth's inner core, which I published 19 in the Proceedings of the Royal Society of London, in 1979. The abstract in its entirety states: 'From observations of nature the suggestion is made that the inner core of the Earth consists not of nickel-iron metal but of nickel silicide'. Whereas Birch had thought that nickel and iron were always alloyed in chondrites, I realized that elemental silicon, found in the metal of *enstatite* chondrites²⁰, under appropriate conditions could cause nickel to precipitate as nickel silicide, an intermetallic compound of nickel and silicon, like the mineral perryite, which had been discovered in enstatite chondrites²¹.

Significantly, a fully crystallized inner core of nickel silicide would constitute a mass virtually identical to the observed mass of the inner core; no such predictability exists for Birch's concept of a partially crystallized nickel–iron metal inner core.

Oxygen rules

Only five elements (Fe, Mg, Si, S and O) constitute about 95% of the mass of each chondrite meteorite and, by implication, about 95% of the mass of each of the terrestrial planets. Four of those elements (F, Mg, Si and S) occur in chondrites in about the same relative proportion as they occur in the outer regions of the Sun, to within a factor of two. Oxygen is the exception, being about 3 times more abundant in the photosphere of the Sun than the sum of the other four elements²². The high relative abundance of oxygen in solar matter poses a serious limitation on the nature of primordial condensates from that medium. The oxidation states of chondritic matter and, by implication, the oxidation states of the terrestrial planets are set by the nature of those condensates and the circumstances of the separation of those condensates from their primordial gases.

The ordinary chondrites comprise about 80% of the meteorites that are observed falling to Earth and, in terms of the five major elements, consist principally of the following minerals: olivine [(Mg, Fe)₂ SiO₄], pyroxene [(Mg, Fe)SiO₃], troilite

[FeS], and metal [Fe]. Considerable confusion has arisen within the Earth sciences by the promulgation of models during the 1970s that incorrectly assumed that this mineral assemblage formed as condensate from a gas of solar composition²³.

Suess and I demonstrated²⁴ from thermodynamic considerations that the oxidized iron content of the silicates of ordinary chondrites was consistent, not with formation from solar matter, but instead with their formation from a gas phase depleted in hydrogen by a factor of about 1000 relative to solar composition. Subsequently, I showed²⁵ that oxygen depletion, relative to solar matter, was also required, otherwise essentially all of the elements would be observed combined with oxygen as they are in the C1 or CI carbonaceous chondrites. Moreover, I showed that if the mineral assemblage characteristic of ordinary chondrites could exist in equilibrium with a gas of solar composition, it is at most only at a single low temperature, if at all. Such a mineral assemblage, therefore, cannot legitimately be assumed to be a primary building component of the Earth. Instead, the ordinary chondrite meteorites appear to have formed from a mixture of two components, re-evaporated after separation from solar gases, one component being an oxidized primitive matter like C1 chondrites, the other being a partially differentiated planetary component from enstatite-chondrite-like matter²⁶.

Even though the scientific underpinnings of the so-called Equilibrium Condensation Model were shown untrue²⁵, the model is still being used as the underlying assumption for other Earth composition models, such as the so-called Bulk Silicate Earth Model. Such models are typically constructed by assuming elemental abundances, usually like those of C1 or CI chondrites, for elements having zero or positive valences, distributing those elements in different regions of the Earth on the basis of various geochemical assumptions, and assigning oxygen on the basis of stoichiometry. But there is no evidence that nature ever obliged those assumptions and it is not scientifically legitimate to use the so-called Equilibrium Condensation Model as justification.

Earth-chondrite relationships

After an inspiring conversation with Lehmann in 1979, I progressed through the following logical exercise: If the inner core is in fact nickel silicide, as I had suggested¹⁹, then the Earth's core must be like the alloy portion of an enstatite chondrite. If the Earth's core is in fact like the alloy portion of an enstatite chondrite, then the Earth's core should be surrounded by a silicate shell like the silicate portion of an enstatite chondrite. This silicate shell, if it exists, should be bounded by a seismic discontinuity, because the silicates of enstatite chondrites have a different and more highly reduced composition than rocks that appear to come from within the Earth's upper mantle. Using the alloy to silicate ratio of the Abee enstatite chondrite and the mass of the Earth's core, by simple ratio proportion, I calculated the mass of that silicate shell. From tabulated mass distributions²⁷, I then found that the radius of that predicted seismic boundary lies within about 1.2% of the radius at the seismic discontinuity that separates the lower mantle from the upper mantle. This logical exercise led me to discover the fundamental quantitative mass ratio relationships connecting the interior parts of the Earth with parts of the Abee enstatite chondrite that are shown in Table 1.

Consider the Earth's core as a percentage by mass of the Earth as a whole, about 32.5%. Similarly, consider the percentage by mass of the alloy portion of each chondrite for which data are available, as shown in Figure 1. Note that, if the Earth has a chondritic composition, as widely believed for good reason, then the Earth is in the main like an enstatite chondrite and not like an ordinary chondrite. An ordinary chondrite-like Earth would not yield a sufficiently massive core. The reason is readily apparent from inspection of the x-axis; the relative oxygen content in chondritic matter primarily determines the relative amount of alloy. The Earth as a whole, and particularly the endo-Earth, the inner 82%, has a state of oxidation like an enstatite chondrite and not like an ordinary chondrite.

The identity of the components of the Abee enstatite chondrite with corresponding components of the Earth means that with reasonable confidence one can understand the composition of the Earth's core by understanding the Abee meteorite or one like it. High pressures, such as prevail within the Earth's core, cannot change the state of oxidation of the core. Not surprisingly, MgSiO₃, the major silicate of the Abee enstatite chondrite, has been

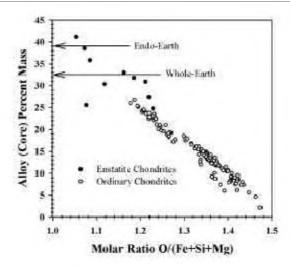


Figure 1. The per cent mass of the alloy component of each of nine enstatite chondrites and 157 ordinary chondrites. The figure shows that, if the Earth is chondritic in composition, the Earth as a whole, and especially the endo-Earth, is like an enstatite chondrite and *not* like an ordinary chondrite. The reason is clear from the abscissa which shows the molar ratio of oxygen to the three major elements with which it combines in enstatite chondrites and in ordinary chondrites. The figure also shows that, if the Earth is chondritic in composition, the Earth as a whole, and especially the endo-Earth, has a state of oxidation like an enstatite chondrite and *not* like an ordinary chondrite. Data from refs 11, 39–41.

Table 1. Fundamental mass ratio comparison between the endo-Earth (core plus lower mantle) and the Abee enstatite chondrite 42

Fundamental Earth ratio	Earth ratio value	Abee ratio value
Lower mantle mass to total core mass	1.49	1.43
Inner core mass to total core mass	0.052	theoretical
		0.052 if Ni₃Si
		0.057 if Ni₂Si
Inner core mass to (lower mantle + core) ma	ss 0.021	0.021

shown to be a stable phase in a perovskite structure at lower mantle pressures^{28–31}.

The oxidation state determines, not only the relative mass of the core, but the elements the core contains. Highly reduced matter, like that of the Abee enstatite chondrite and the endo-Earth (i.e. the core and lower mantle), was separated from solar gases under conditions that severely limited the oxygen content³². As a consequence certain elements, including Si, Mg, Ca, Ti, U and Th, which would occur entirely as silicate-oxides in ordinary chondrites, occur in part in the alloy portion of the Abee enstatite chondrite and in the Earth's core. Being unable to form oxides, those core-elements compete on the basis of chemical activity and may be accommodated otherwise, for example, as sulfides. Commercially, to desulfurate steel, calcium (Ca) or magnesium (Mg) is intentionally introduced into the molten alloy to form CaS and MgS, which precipitates at a high temperature and floats to the surface. The expectation within the Earth's core is for CaS and MgS to precipitate at a high temperature and to float to the top of the core; I have suggested that the 'islands' of matter at the core mantle boundary³³ consist of such CaS and MgS^{4,14,34}.

Whereas the gross features of the endo-Earth appear relatively simple, consistent with the identification of that part being like an enstatite chondrite, the upper mantle displays several seismic discontinuities suggestive of different layers. The challenge is not to make models assuming their compositions but, rather, to identify with certainty the compositions of those layers by discovering fundamental quantitative relationships. If the Earth is chondritic in composition, the upper mantle may be expected to consist mainly of mixtures of the components from the two 'primitive' chondritic matter formation reservoirs that yielded the highly oxidized matter like the C1 or CI chondrites and the highly reduced matter like the enstatite chondrites.

The tasks ahead

Because of the fundamental mass ratio relationships connecting the Earth's core and lower mantle to corresponding parts of the Abee enstatite chondrite (Table 1 and Figure 1), one may use analytical data for that meteorite, or one like it, to understand the nature of deep-Earth chemistry, particularly the partitioning of elements between the core and lower mantle, which is governed by oxygen fugacity and sulfur fugacity. Regrettably, the Earth science community has systematically failed to exploit a real opportunity to advance scientific understanding. The best and most comprehensive work on the mineralogy and chemical relationships among enstatite chondrites35 was published in 1968, at a time when the electron microprobe was not yet fully perfected. The best data on the distribution actinide elements among components of the Abee meteorite³⁶ was published in 1982. Imagine the benefit to be realized from precise and thorough investigations of enstatite chondrites using current, state-of-the-art technology.

Vast resources have been expended for high-pressure diamond-anvil laboratory experiments. Regrettably the results to date are of little relevance for the Earth's interior, being focused mainly on the phase relationships of iron metal. What is needed is to elucidate the phase relations in the system Fe-Ni-S-Si-Mg-Ca at core pressures and temperatures as discussed by Herndon³⁷. In particular, it is important to reveal the p-T conditions under which nickel silicide will precipitate, as these data may place constraints on the temperature at the outer boundary of the inner core. It is important to determine the range of possible compositions of nickel silicide and the physical properties of each

From the standpoint of geo-antineutrino investigations, it is vital to determine the partitioning of naturally occurring radioactive elements between the Earth's core and mantle in the system Fe-Si-Mg-S-O for compositions near that of the Abee enstatite chondrite. Such investigations should go hand-in-hand with state-of-the-art mineralogical and chemical investigations of the enstatite chondrites. Simply showing, for example, that an element

like potassium reacts with iron at core p-T is without meaning, as the oxygen fugacity and sulfur fugacity are dominant considerations. As an aid to investigators planning geo-antineutrino projects, Herndon and Edgerley³⁸ recently set forth preliminary estimates of the abundances and distribution of fissionable and nonfissionable radionuclides within the deep interior of the Earth. Ultimately, these estimates should be revised when modern, state-of-the-art data become available.

- Raghavan, R. S., 2002, arXiv: hep-ex/ 0208038 Aug 2002.
- de Meijer, R. J., van der Graaf, E. R. and Jungmann, K. P., *Nucl. Phys. News Int.*, 2004, 14, 20–25.
- Domogatski, G., Kopeikin, L., Mikaelyan, L. and Sinev, V., 2004, arXiv: hepph/0401221 v1 28 Jan 2004.
- 4. Herndon, J. M., *J. Geomag. Geoelectr.*, 1993, **45**, 423–437.
- 5. Herndon, J. M., *Proc. Natl. Acad. Sci. USA*, 2003, **100**, 3047–3050.
- Hollenbach, D. F. and Herndon, J. M., Proc. Natl. Acad. Sci. USA, 2001, 98, 11085–11090.
- 7. Wiechert, E., Nachr. K. Ges. Wiss. Goettingen, Math.-Kl., 1897, 221–243.
- 8. Cavendish, H., *Philos. Trans. R. Soc. London*, 1798, **88**, 469–479.
- 9. Oldham, R. D., Q. T. Geol. Soc. London, 1906, **62**, 456–476.
- 10. Jeffreys, H., *The Earth*, Cambridge, 1929.

- 11. Jarosewich, E., *Meteoritics*, 1990, **25**, 323–337.
- 12. Wiik, H. B., Commentat. Phys.-Math., 1969, **34**, 135–145.
- 13. Urey, H. C. and Craig, H., *Geochim. Cosmochim. Acta*, 1953, 4, 36–82.
- Herndon, J. M., Proc. Natl. Acad. Sci. USA, 1996, 93, 646–648.
- 15. Mason, B., *Am. Mus. Novit.*, 1962, **2085**, 1–20.
- Clayton, R. N., Annu. Rev. Earth Planet. Sci., 1993, 21, 115–149.
- Lehmann, I., Publ. Int. Geod. Geophys. Union, Assoc. Seismol., Ser. A, Trav. Sci., 1936, 14, 87–115.
- 18. Birch, F., Am. J. Sci., 1940, 238, 192-211.
- Herndon, J. M., Proc. R. Soc. London, 1979, A368, 495–500.
- 20. Ringwood, A. E., *Geochim. Cosmochim. Acta*, 1961, **25**, 1–13.
- 21. Ramdohr, P., *Abh. D. Akad. Wiss. Ber.*, *Kl. Chem.*, *Geol.*, *Biol.*, 1964, **5**, 1–20.
- Anders, E. and Grevesse, N., Geochim. Cosmochim. Acta, 1989, 53, 197–214.
- 23. Larimer, J. W. and Anders, E., *Geochim. Cosmochim. Acta*, 1970, **34**, 367–387.
- Herndon, J. M. and Suess, H. E., Geochim. Cosmochim. Acta, 1977, 41, 233–236.
- 25. Herndon, J. M., *Proc. R. Soc. London*, 1978, **A363**, 283–288.
- Herndon, J. M., 2004, arXiv: astro-ph/ 0405298 15 May 2004.
- Dziewonski, A. M. and Gilbert, F., Geophys. J. R. Astr. Soc., 1972, 72, 393–446.
- 28. Liu, L., Earth Planet. Sci. Lett., 1976, 31, 200–208.
- Ito, E. and Matsui, Y., Earth Planet. Sci. Lett., 1978, 38, 443–450.

- 30. Chaplot, S. L., Choudhury, N. and Rao, K. R., *Am. Min.*, 1998, **83**, 937–941.
- 31. Chaplot, S. L. and Choudhury, N., Am. Min., 2001, **86**, 752–761.
- 32. Herndon, J. M. and Suess, H. E., *Geochim. Cosmochim. Acta*, 1976, **40**, 395–399.
- 33. Vidale, J. E. and Benz, H. M., *Nature*, 1993, **361**, 529–532.
- 34. Herndon, J. M., *Phys. Earth Plan. Inter.*, 1998, **105**, 1–4.
- 35. Keil, K., J. Geophys. Res., 1968, 73, 6945–6976.
- Murrell, M. T. and Burnett, D. S., Geochim. Cosmochim. Acta, 1982, 46, 2453– 2460.
- 37. Herndon, J. M., Eos, Trans. Am. Geophys. Union, 1998, **79**, 451–456.
- 38. Herndon, J. M. and Edgerley, D. A., 2005, arXiv:hep-ph/0501216 24 Jan 2005.
- Baedecker, P. A. and Wasson, J. T., Geochim. Cosmochim. Acta, 1975, 39, 735– 765.
- Kallemeyn, G. W. and Wasson, J. T., Geochim. Cosmochim. Acta, 1981, 45, 1217–1230.
- Kallemeyn, G. W., Rubin, A. E., Wang,
 D. and Wasson, J. T., Geochim. Cosmochim. Acta, 1989, 53, 2747–2767.
- 42. Herndon, J. M., *Proc. R. Soc. London*, 1980, **A372**, 149–154.

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